

Late Jurassic-Early Cretaceous kimberlite, carbonatite and ultramafic lamprophyric sill and dyke swarms from the Bomethra area, northeastern Oman

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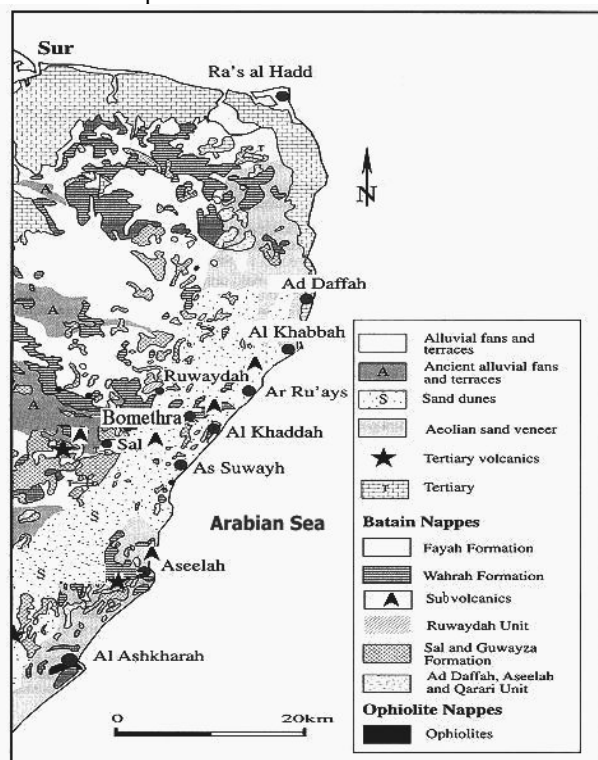
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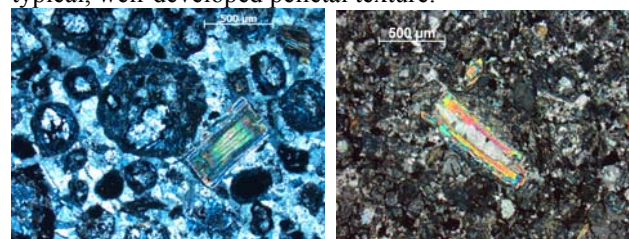
The Batain Nappes in Oman is represented by an allochthonous unit mainly consists of sedimentary rocks with subordinate volcanic rocks and minor intrusions, deposited in the Batain Basin along the eastern Oman continental margin during the Permian to the Late Cretaceous times. The allochthonous units in the Batain Plain were emplaced in a NW direction on to the Arabian plate at the Cretaceous/Tertiary boundary. The Nappes contain ophiolite, ultramafic lamprophyres, carbonatite and kimberlites occurring as pipes in the Asseelah area and as plugs, sills and dykes of several meters thick and about 6 km long in the Bomethra area. The majorities of these rocks are hosted by the Late Jurassic to Early Cretaceous Wahra Formation, which comprises a series of radiolarian cherts, shales and clay stones, and are found over the whole Batain plain.



Petrography and Mineralogy

The dyke and sill swarms are located approximately 20 km northeast of the Aseelah pipe (Nasir et al., 2008) in the Bomethra area (Map above). The main sill has a thickness of 1-2 m and extends to 6 km in N40°E direction. Two blows, 200x500 m exposure of effusive rocks occur within the main sill. They are composed of lapilli and bombs, mixed with pieces of radiolarian cherts.

Two varieties of ultramafic rocks have been observed: one is medium- to coarse-grained mica-carbonate-rich damtjernite with abundant calcite-ocelli and/or pelletal lapilli, and the other is massive fine-grained tuffitic aillikite (Figures below). Microscopically, the damtjernite exhibits a range of textures from well-developed segregations of calcite and phlogopite with abundant microlitic apatite and phlogopite, to the more typical, well-developed pelletal texture.

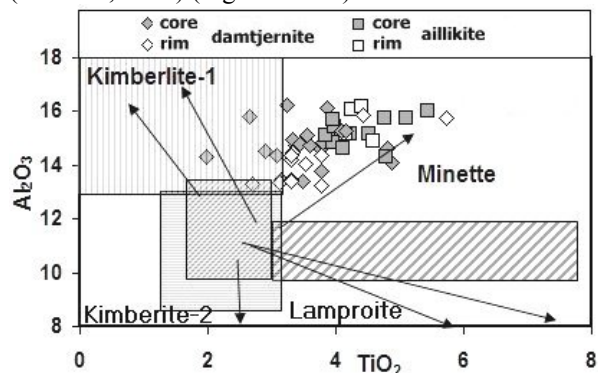


The groundmass in both varieties is composed of mainly of calcite, phlogopite, and apatite. Cr-Spinel, magnetite and rutile are common. Orthoclase, diopside and richterite occur mainly in the damtjernite.

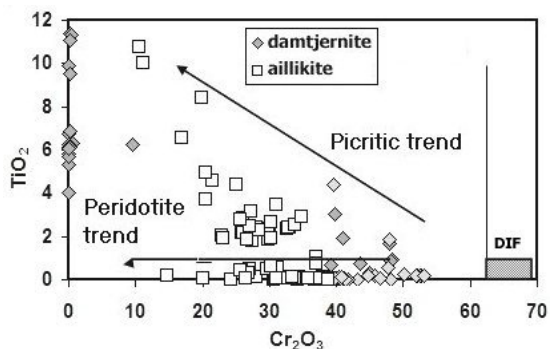
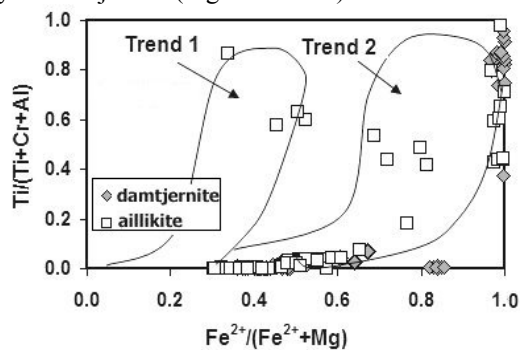
Carbonate typically occurs together with phlogopite and apatite as fine-grained (<0.1 mm) aggregates of xenomorphic grains in the ground mass. Primary Sr-bearing calcite (0.1 to 0.15 wt%) is the dominant carbonate phase and is commonly present in the groundmass of all samples.

Apatite is a late-crystallizing groundmass phase. The grains are euhedral, prismatic with hexagonal base sections. Abundant apatite occurs as acicular grains and as larger more prismatic grains grown primarily within calcite segregations. They are relatively Si-rich (0.9-1.8 wt% SiO₂) and Sr-poor (<1.5 wt% SrO).

Phlogopite is a major constituent in the Bomethra ultramafic rocks, generally forming > 30 vol. % of the mineral assemblage. It occurs as euhedral six-sided macrocryst, phenocrysts, and unehedral micro-phenocrysts and as small groundmass crystals characterized by subhedral to euhedral laths (0.05 to 0.1 mm). The phlogopite is titania aluminous and plots along a trend described for minette and alnöite rocks (Mitchell, 1986) (Figure below).



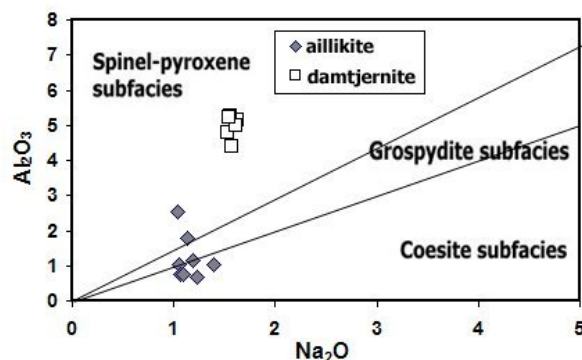
Spinel is a ubiquitous mineral phase in all samples. It occurs as macrocrysts, primary groundmass minerals, and occasionally in association with ilmenite and rutile. Small (<0.2 mm) euhedral-to-subhedral spinels comprise a significant portion of the groundmass mineral assemblage of both aillikite and damtjernite. TiO₂ contents are typically <1 wt.%, increasing with evolution to as high as ≈11.5 wt.%. Two separate evolutionary trends of primary magmatic groundmass spinels are found: magnesian ulvöspinel magmatic trend (trend 1) and titanomagnetite magmatic trend (trend 2). The former occurs in the aillikite groundmass spinels, while the latter is characteristic only of damtjernite (Figures below).



One forsteritic olivine grain and one G4 garnet grain, beside abundant chromite grains were recovered from the heavy mineral concentrates. Garnet grain has a size

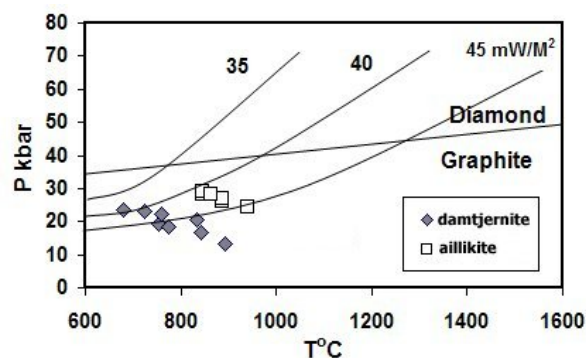
of 0.43 x 0.35 mm. It is a spalled, oval-flattened, pale yellow-orange grain. Pure orthoclase occurs only in the damtjernite.

Cr-diopside and richterite occur sporadically in the carbonate groundmass. Diopside from the damtjernite has lower Al₂O₃ and Na₂O than those from the aillikite (Figure below). Rutile, chlorite and serpentine are common accessory minerals.



P-T estimation

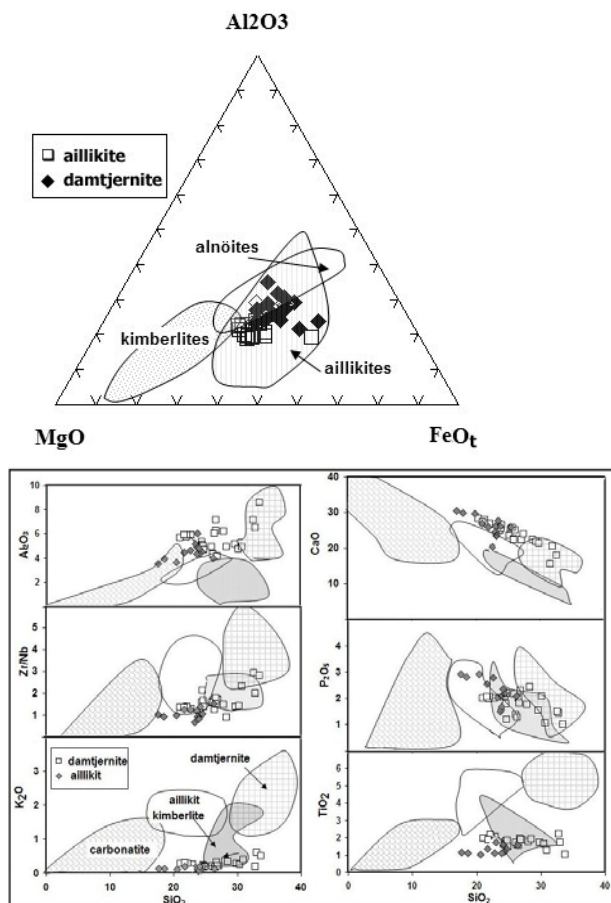
The crystallization pressure of clinopyroxene as estimated by the clinopyroxene barometer of Nimis (1995) can be bracketed between 1.3 and 2.4 GPa for the damtjernite and between 24-29 GPa for the aillikite. Temperatures vary between 680 and 890°C for the damtjernite and between 840-940°C for the aillikite, (Figure below).



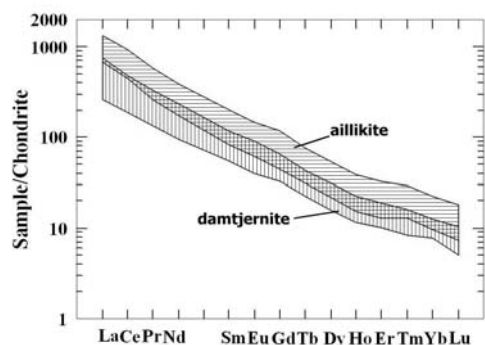
Whole Rock Geochemistry

The Bomethra rocks are strongly silica undersaturated (17-34 wt. % SiO₂), with high Al₂O₃ (3.5-9.5 wt.%) and variable TiO₂ (1-3 wt.%). They have low MgO (5-11 wt. %), which never approach MgO-values similar to kimberlites. The Bomethra rocks have petrographic and geochemical affinities with aillikite. On the FeO_T-MgO-Al₂O₃ ternary discriminant plot, and SiO₂ vs. selected major elements (Figures left below), these rocks generally fall in the aillikite field.

The incompatible trace element abundances show a significant enrichment in most elements, particularly Nb, Ta and LREEs and pronounced troughs at Rb, K, Pb, P, Hf, and Ti (Figure left below).



The REE patterns are all remarkably similar and show strong LREE enrichment with La_n/Yb_n between 30-70 (Figure below).



K-Ar whole rock age dating yielded an age between 130-140 Ma, similar to that of the Asseelah aillikites and carbonatites (Nasir et al., 2008). The initial Sr, Nd and Pb isotope compositions of the aillikite and damtjernite have strong affinity with ultramafic lamprophyres and are similar to the Reunion mantle plume

Discussion

The Bomethra rocks show very low SiO₂ and MgO contents, much lower than most common alkaline magmas. They have petrographic and geochemical affinities with aillikite, damtjernite, and kimberlite, and resemble experimentally produced melts of synthetic carbonated garnet peridotite in the simple CMAS-CO₂ system close to 3 GPa. The high abundance of REE and most incompatible elements may be interpreted as a result of derivation from a lower degree of partial melts or due to enrichment in the mantle source. The K and Ti depletions may be related to a residual phlogopite phase in the source mantle during partial melting. The abundance of magmatic carbonate suggests that the primary magmas were derived from a carbonated mantle. The initial Sm-Nd and Pb isotope are close to the value of the Reunion mantle plume.

There is no obvious age progression for the various types of magmatic activity in the Batain Nappes. It is quite possible, therefore, that all of the magmatic rocks seen in the Batain plain originated in a single event, triggered by the arrival of upwelling mantle plume beneath the metasomatised Indo-Arabian lithosphere. The influx of heat caused melting of the carbonated garnet peridotite asthenospheric mantle forming carbonatite, aillikite, damtjernite and other more fractionated alkaline magmas. The Late Jurassic-Early Cretaceous magmatism in eastern Oman is related spatially and compositionally to mantle upwelling associated with the Reunion mantle plume.

References

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